



WAXI - West African Exploration Initiative

IXOA - L'Initiative d'Exploration Ouest Africaine

Centre for **EXPLORATION**
TARGETING



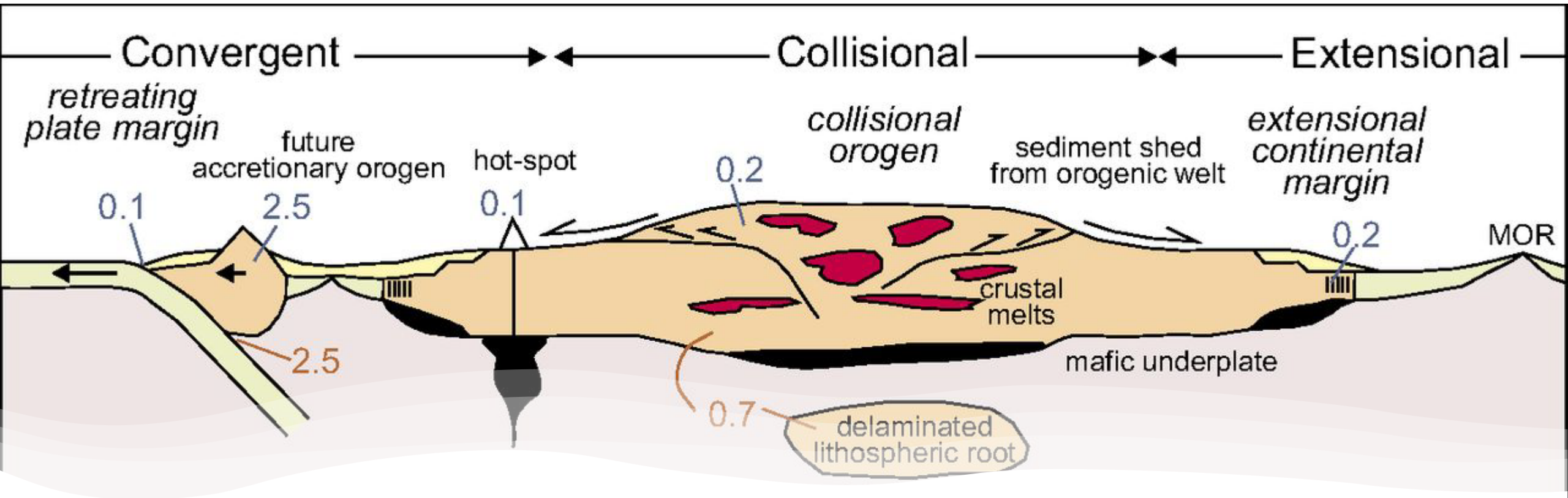
Radiogenic Isotopes as tool to peer into the crustal growth, tectonic evolution and metal fertility of continental masses

Nicolas Thébaud, Quentin Masurel, Luis Para-Avila, Aurélien Eglinger, Julien Perret, Augustin Yao Koffi, Ousmane Wane, Helen McFarlane, Graham Begg, Guillaume Duclaux, Mark Jessell, Olivier Vanderhaeghe, Lenka Baratoux, Mahamadou Diallo, Alain Kouamelan, Tony Kemp, Noreen J. Evans and Armin Zeh



THE UNIVERSITY OF
WESTERN
AUSTRALIA

IASTG - conference Accra July 2026



Cawood et al., 2013

Where does the continental crust forms?

Mechanism for generating TTGs

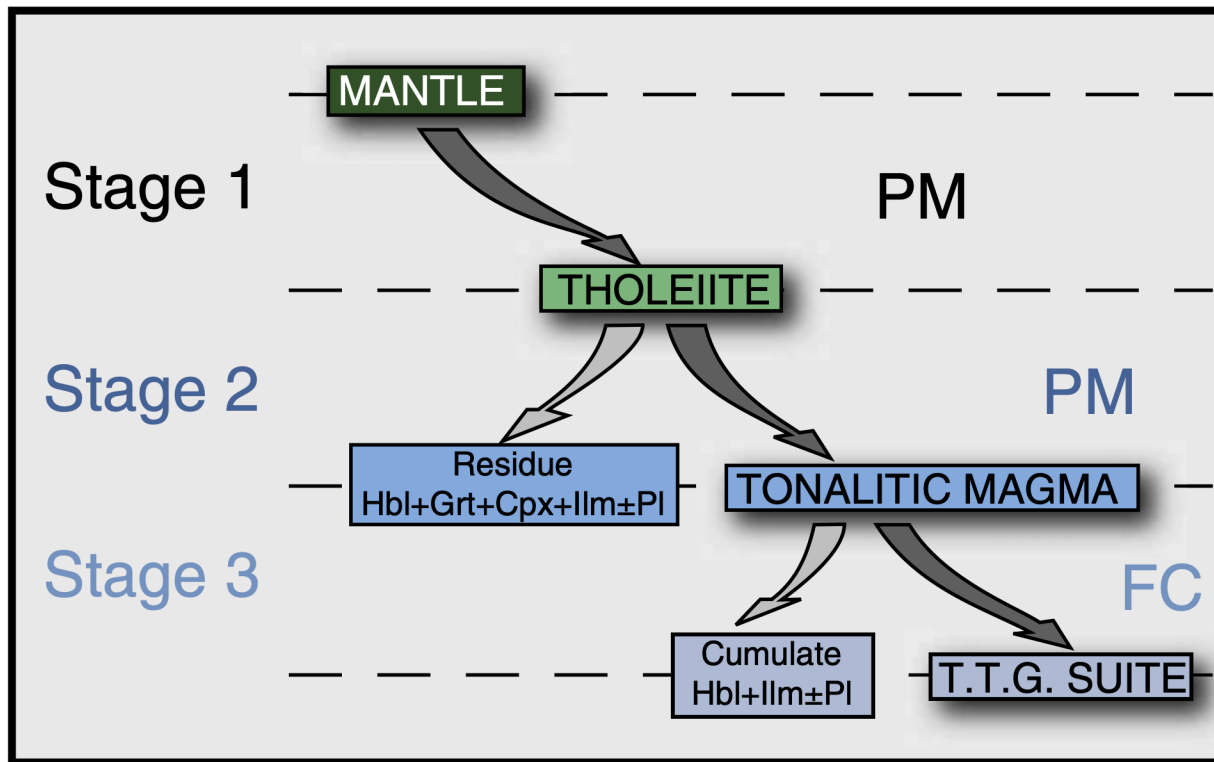
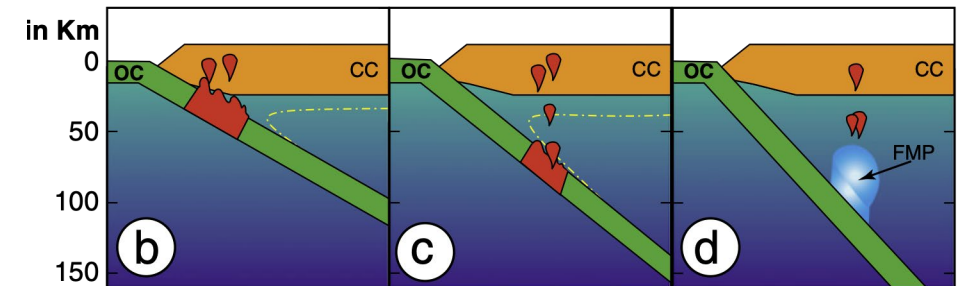
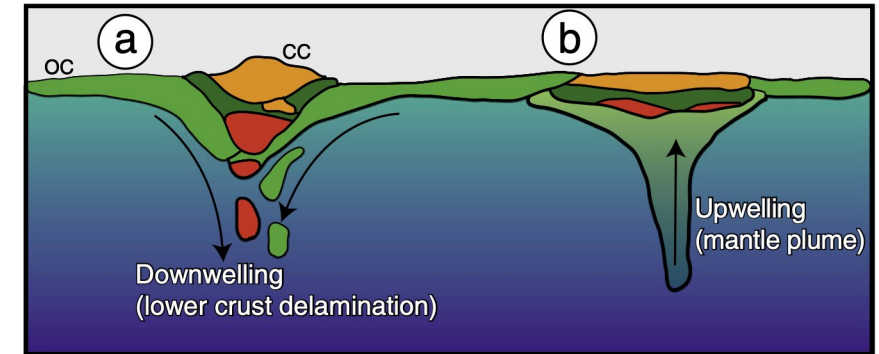


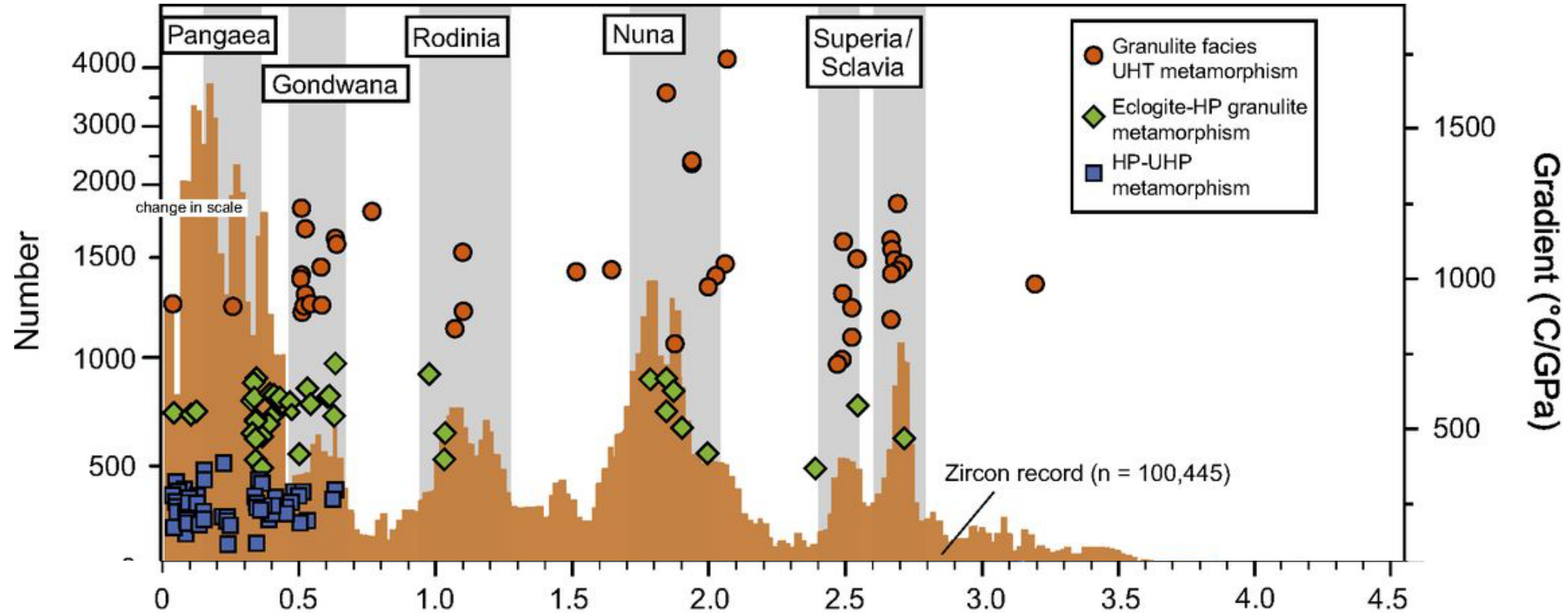
Fig. 8. Schematic diagram summarising the succession of the different mechanisms implied in Archaean TTG genesis (after [Martin, 1993](#)). Hbl = hornblende; Grt = garnet; Pl = plagioclase; Cpx = clinopyroxene; Ilm = ilmenite; PM = partial melting; FC = fractional crystallisation.

without or with plate tectonics ?



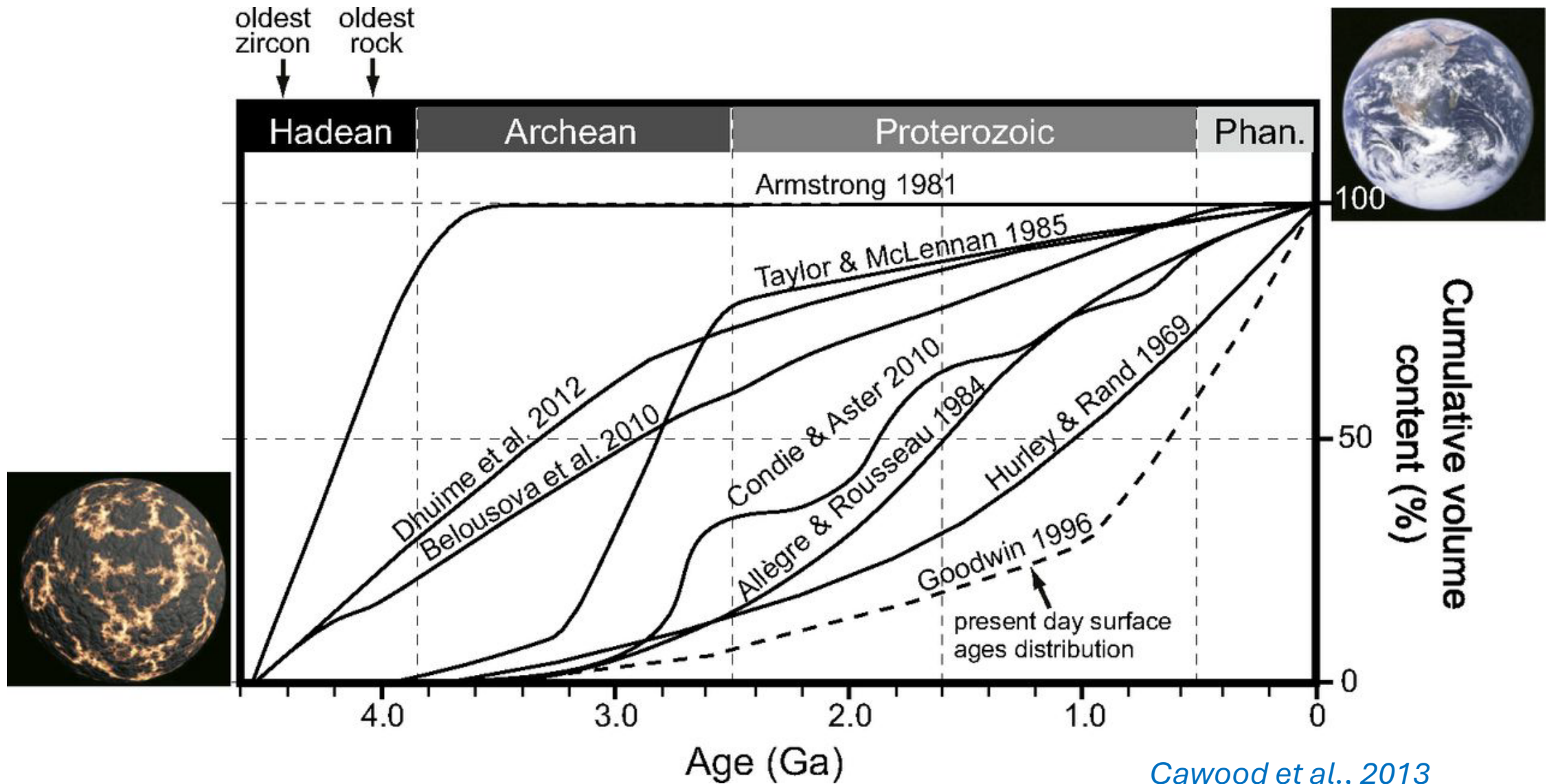
Moyen and Martin., 2012

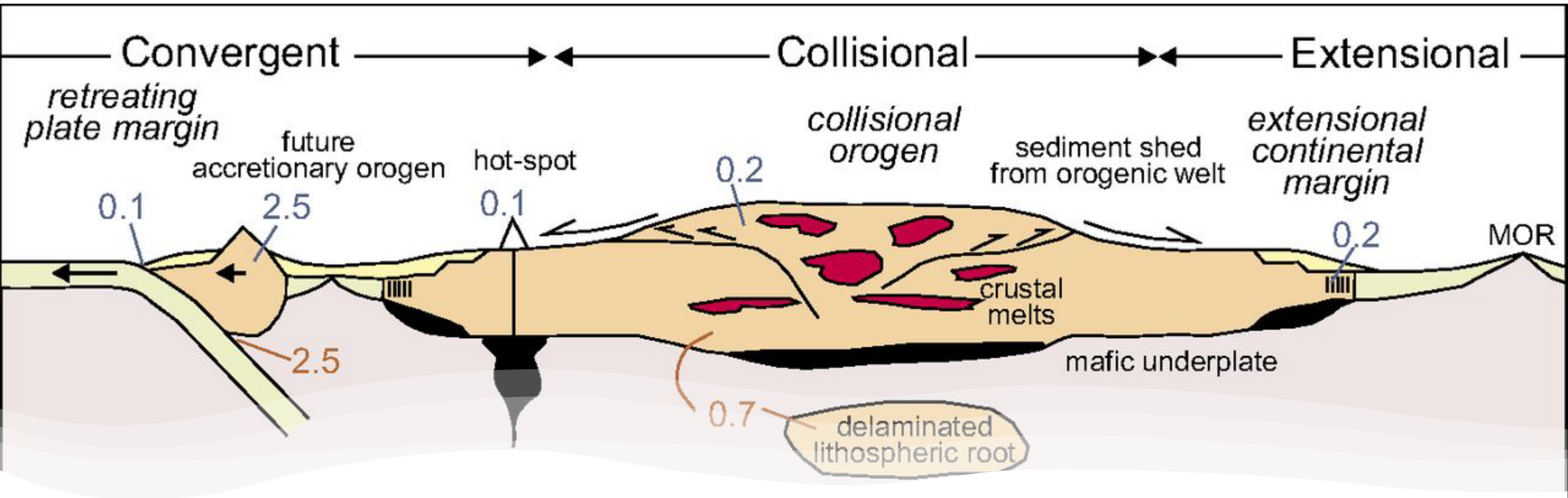
Emergence of the felsic crust ?



Histogram of over 100,000 detrital zircon analyses showing several peaks in their U-Pb crystallization ages over the course of Earth history (Voice et al., 2011 and modified after Cawood et al., 2013)

Emergence of the felsic crust ?

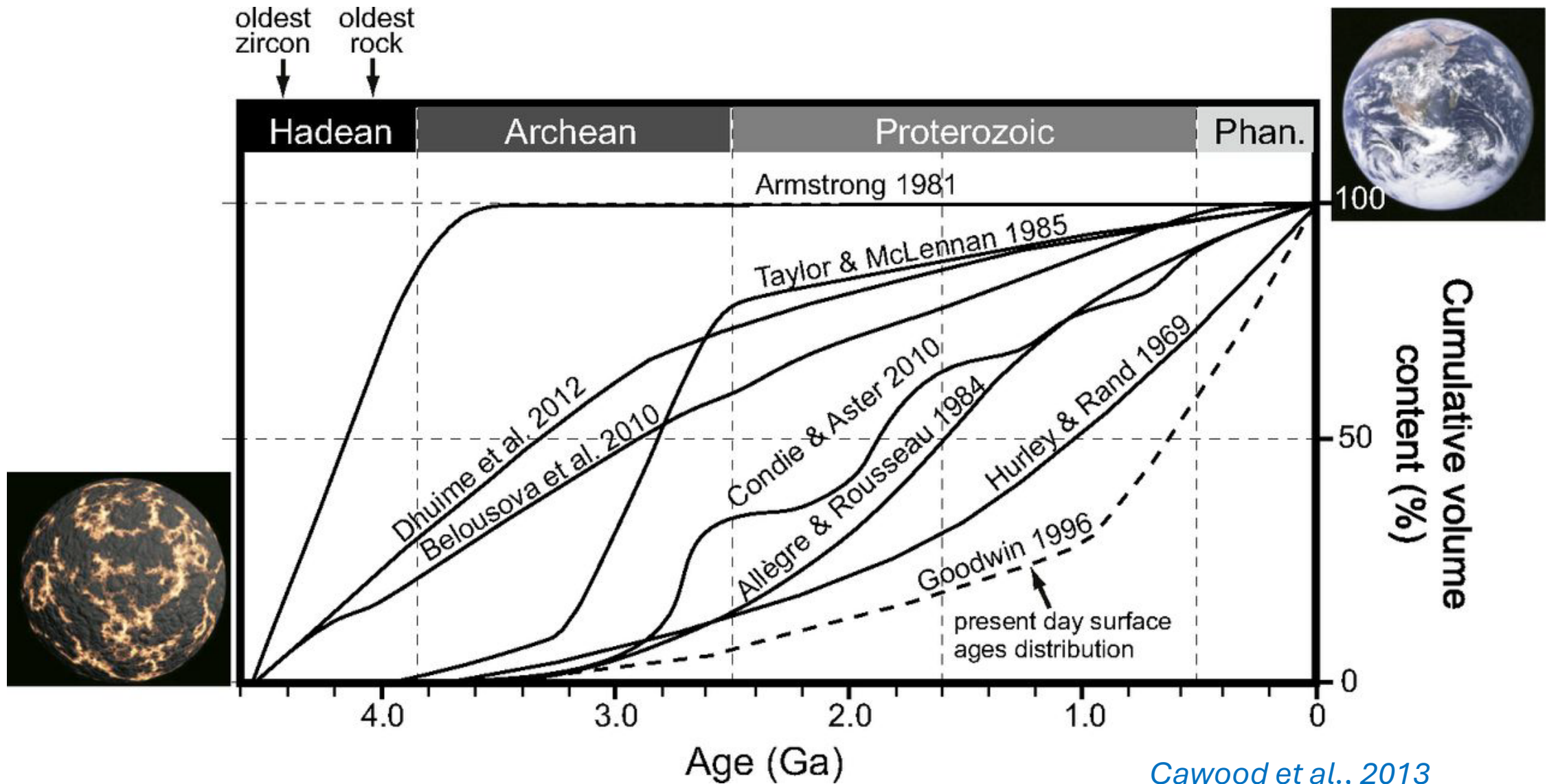




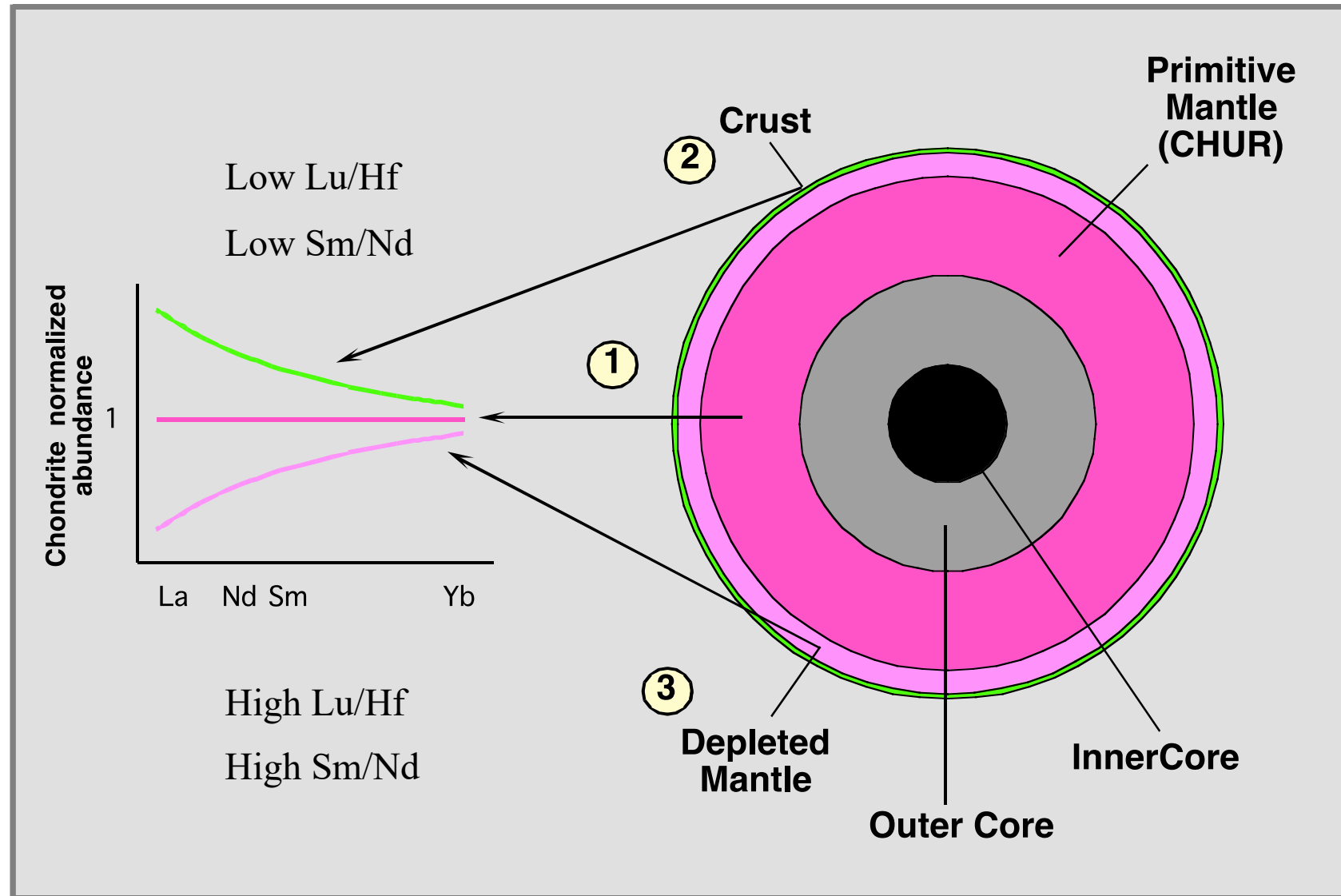
Cawood et al., 2013

Where does the continental crust disappear?

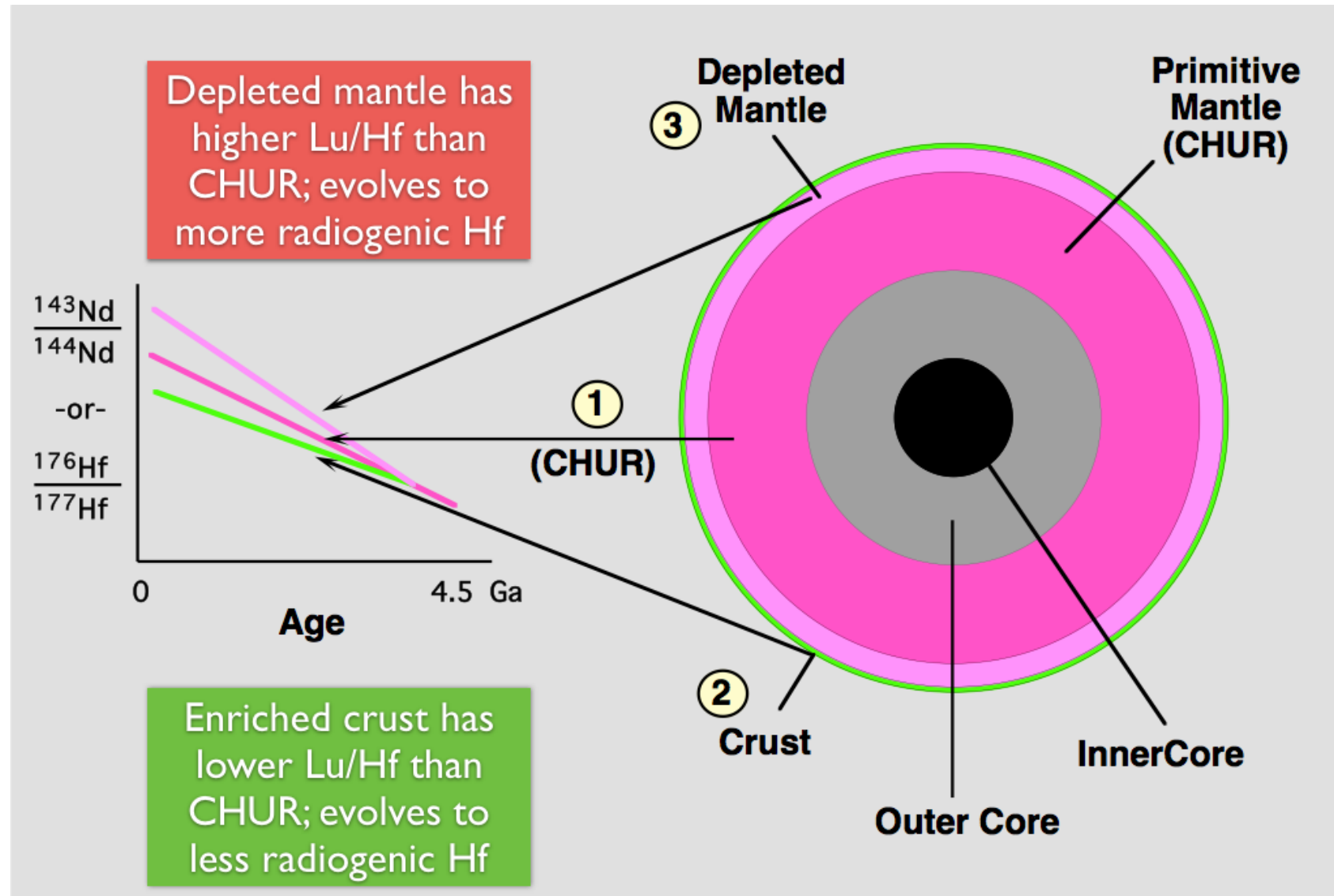
Emergence of the felsic crust ?



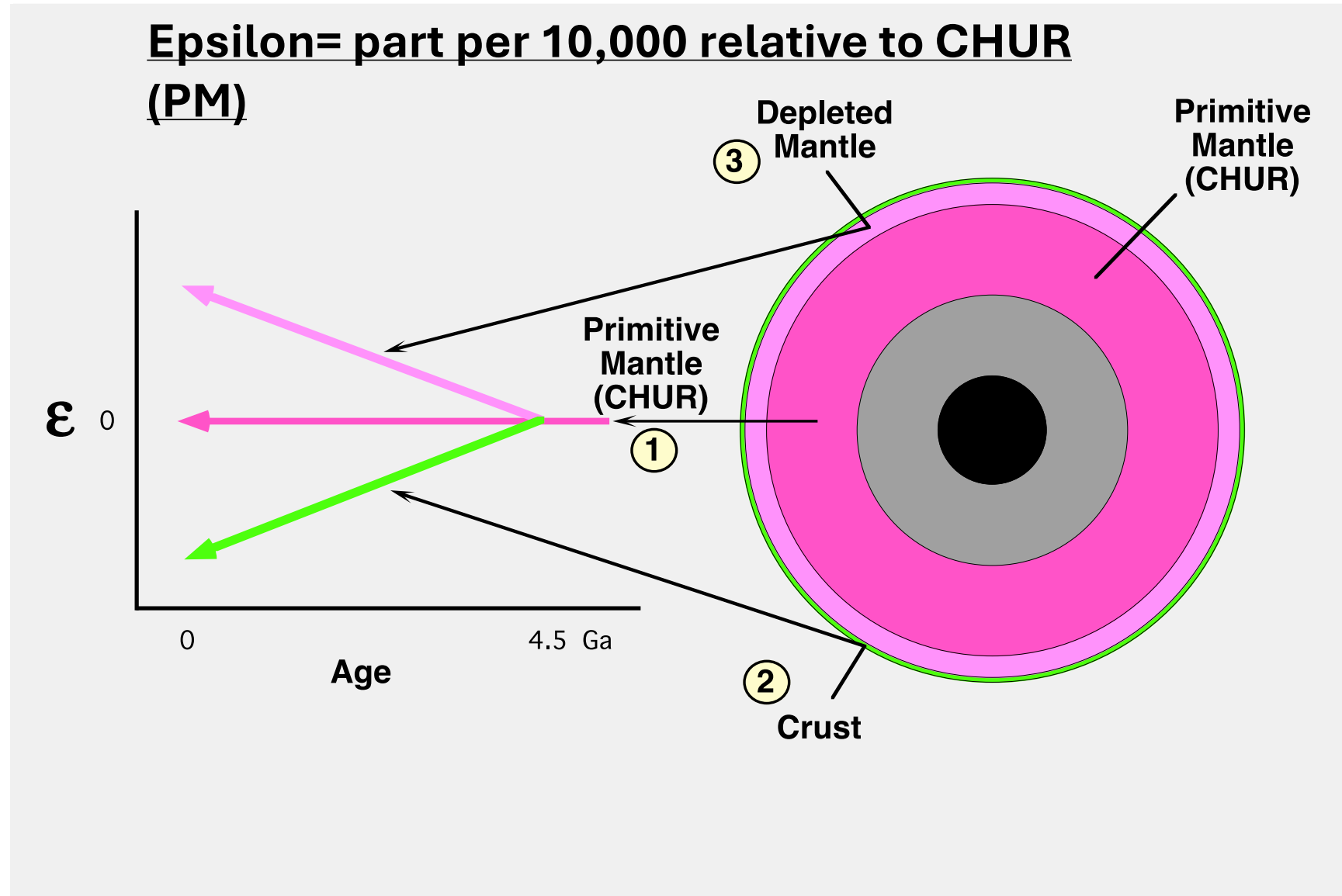
Peering further into the crustal growth using radiogenic isotopes Lu/Hf and Nd/Sm systematic

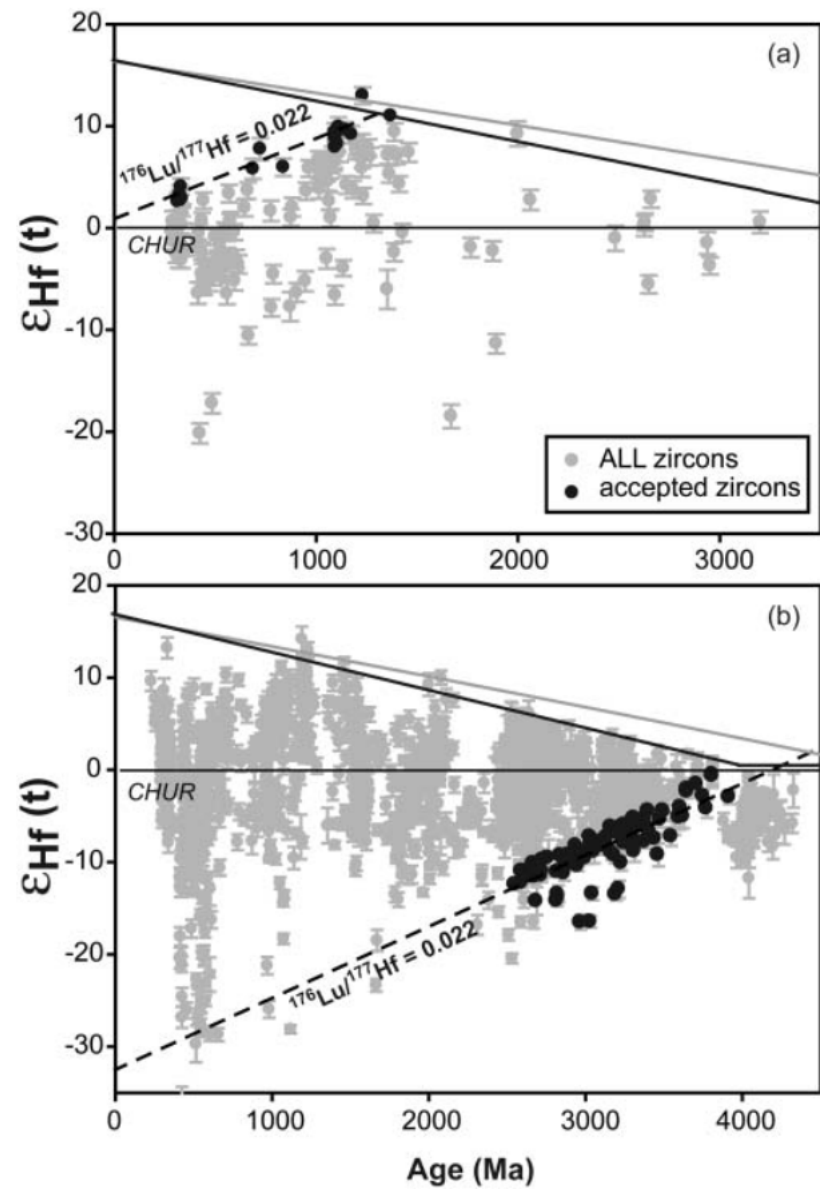
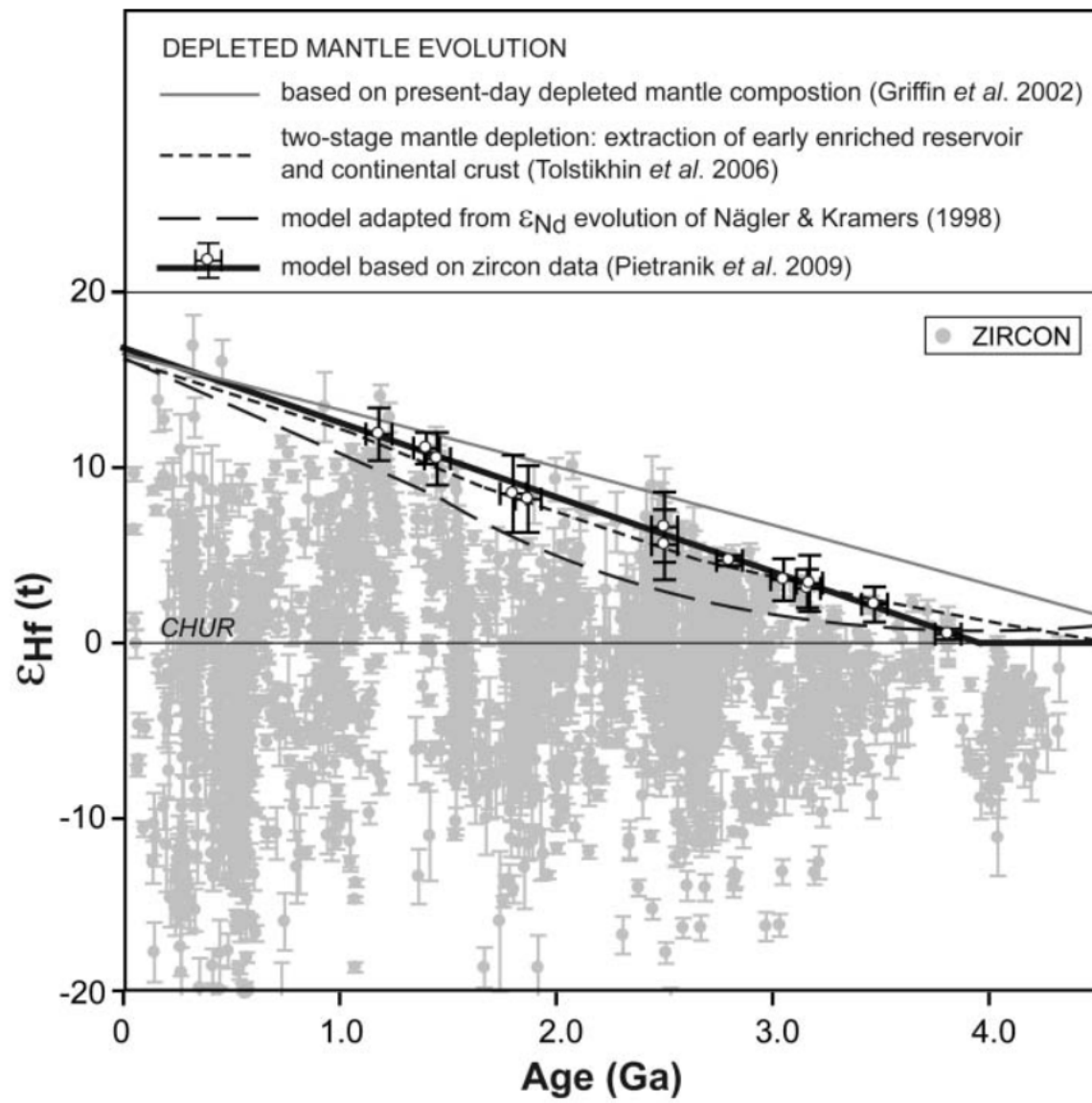


Peering further into the crustal growth using radiogenic isotopes Lu/Hf and Nd/Sm systematic

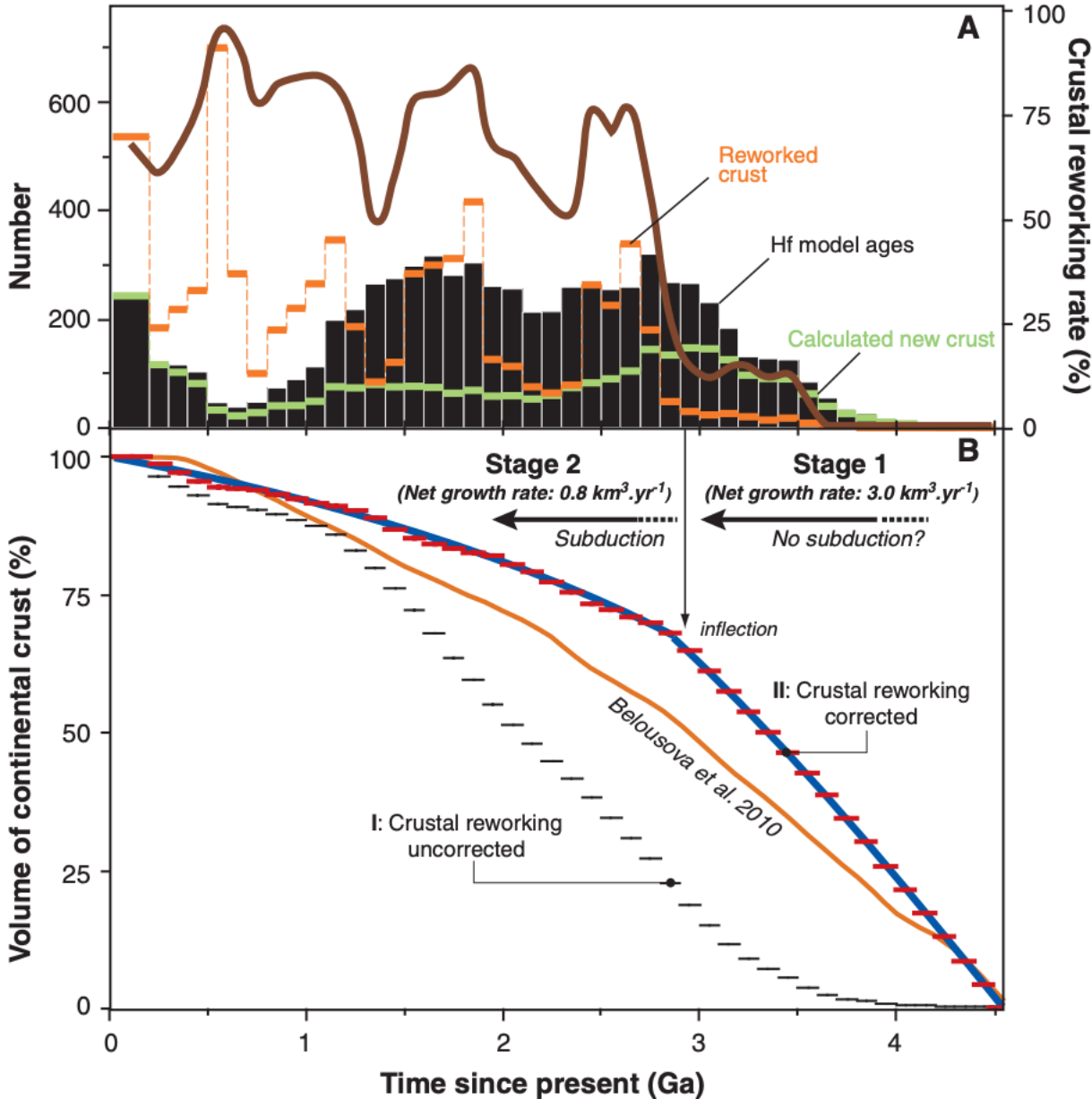


Peering further into the crustal growth using radiogenic isotopes Lu/Hf and Nd/Sm systematic





Emergence and preservation of the felsic crust ?



Why does the ancient crust best preserved?



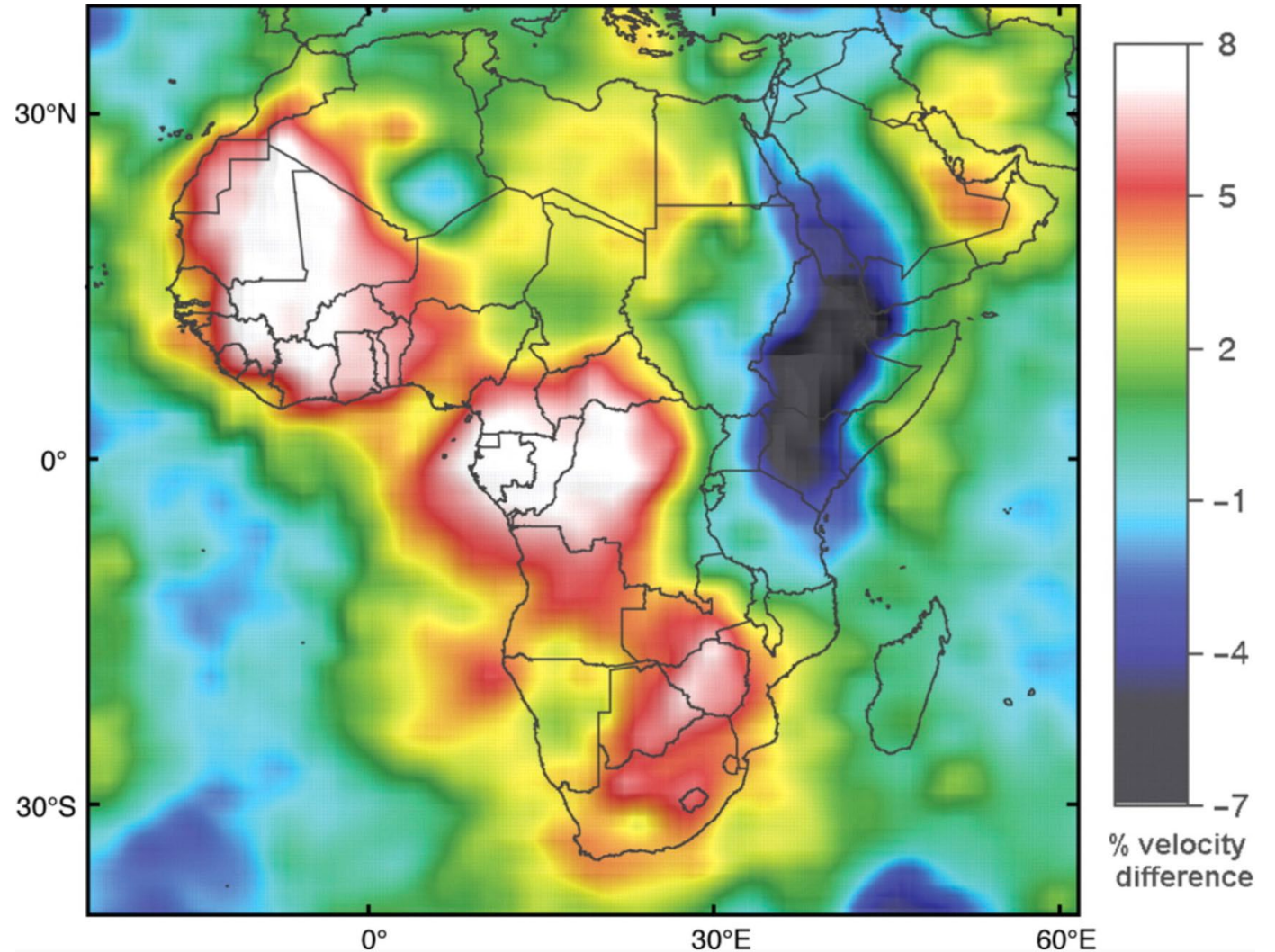
Role of the continental lithospheric mantle?

Credit Griffin



Africa 100-175 km Vs tomography

- Red/white = large Vs
- blue = small Vs



Source data Deen et al. 2006, published in Griffin et al., 2009

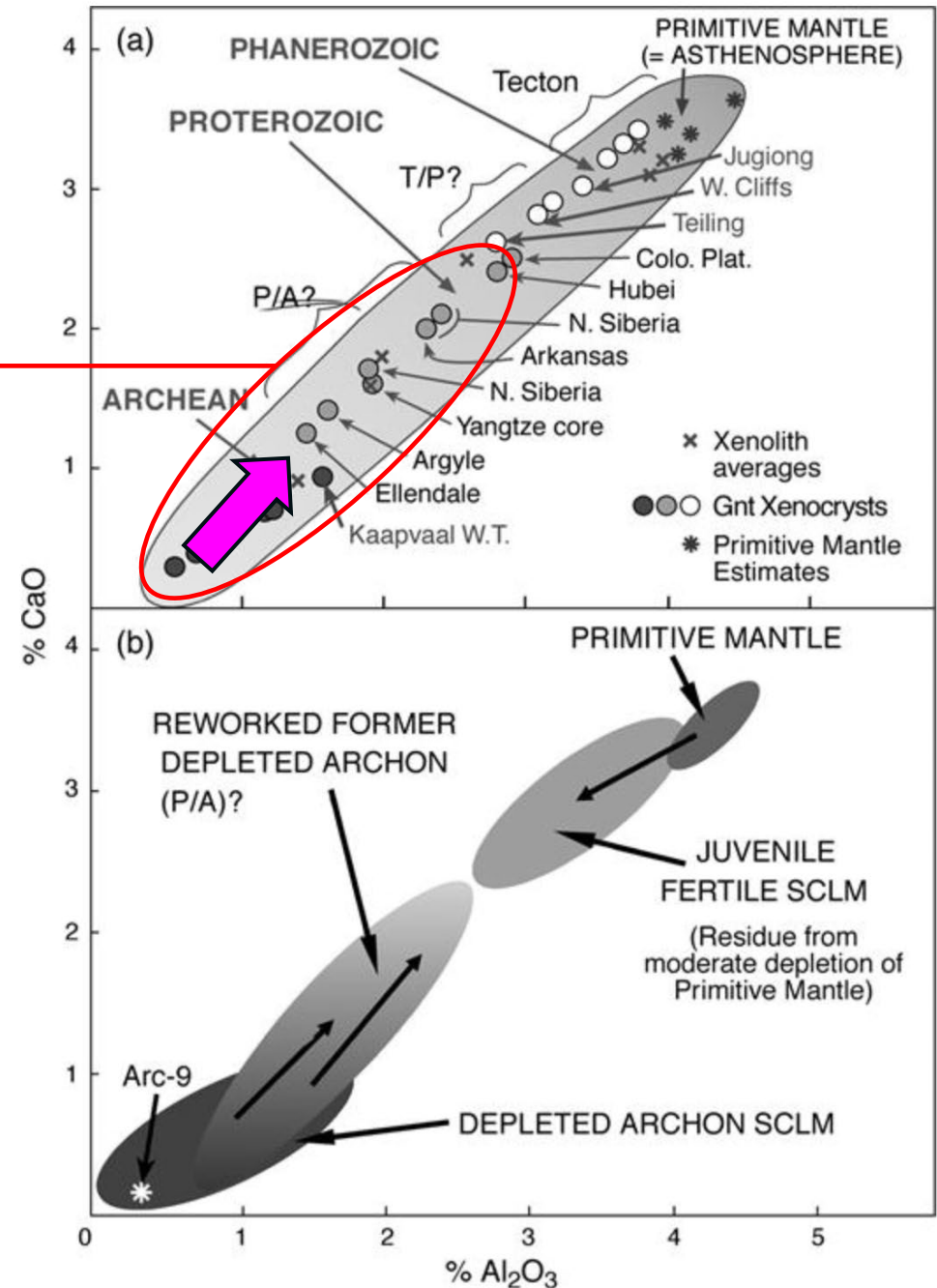
Contributions from mantle xenoliths

Archean CLM!

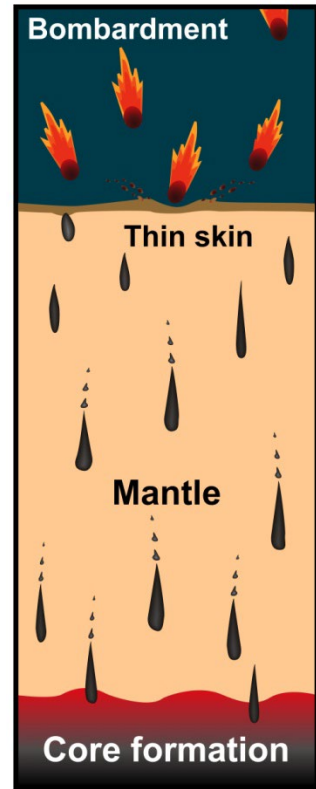
PLUME RESIDUE

- dominated by **Mg-rich** (fosterite) Olivine
- Unique to melting of high T Archean plumes
- Distinctive **High Velocity**

Griffin et al., 2009

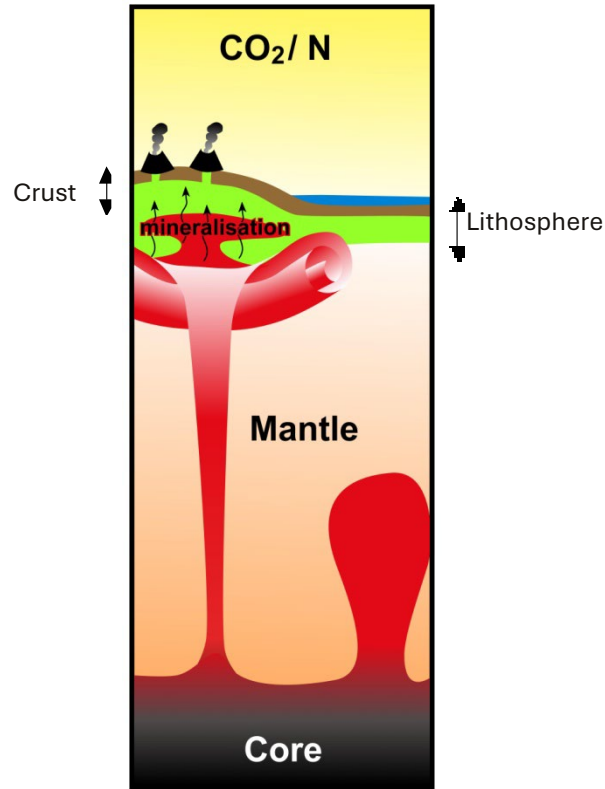


CLM – Archean Genesis



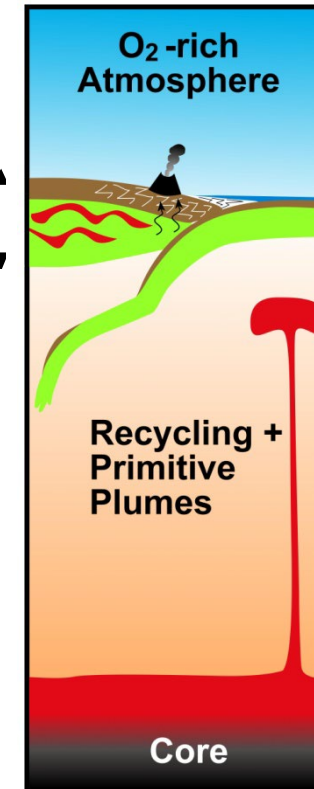
Hadean

- no stable CLM
- no preservation of the crust



Archean (3.6-3.0 Ga)?

- mantle overturns
- formation of stable CLM
- subduction?

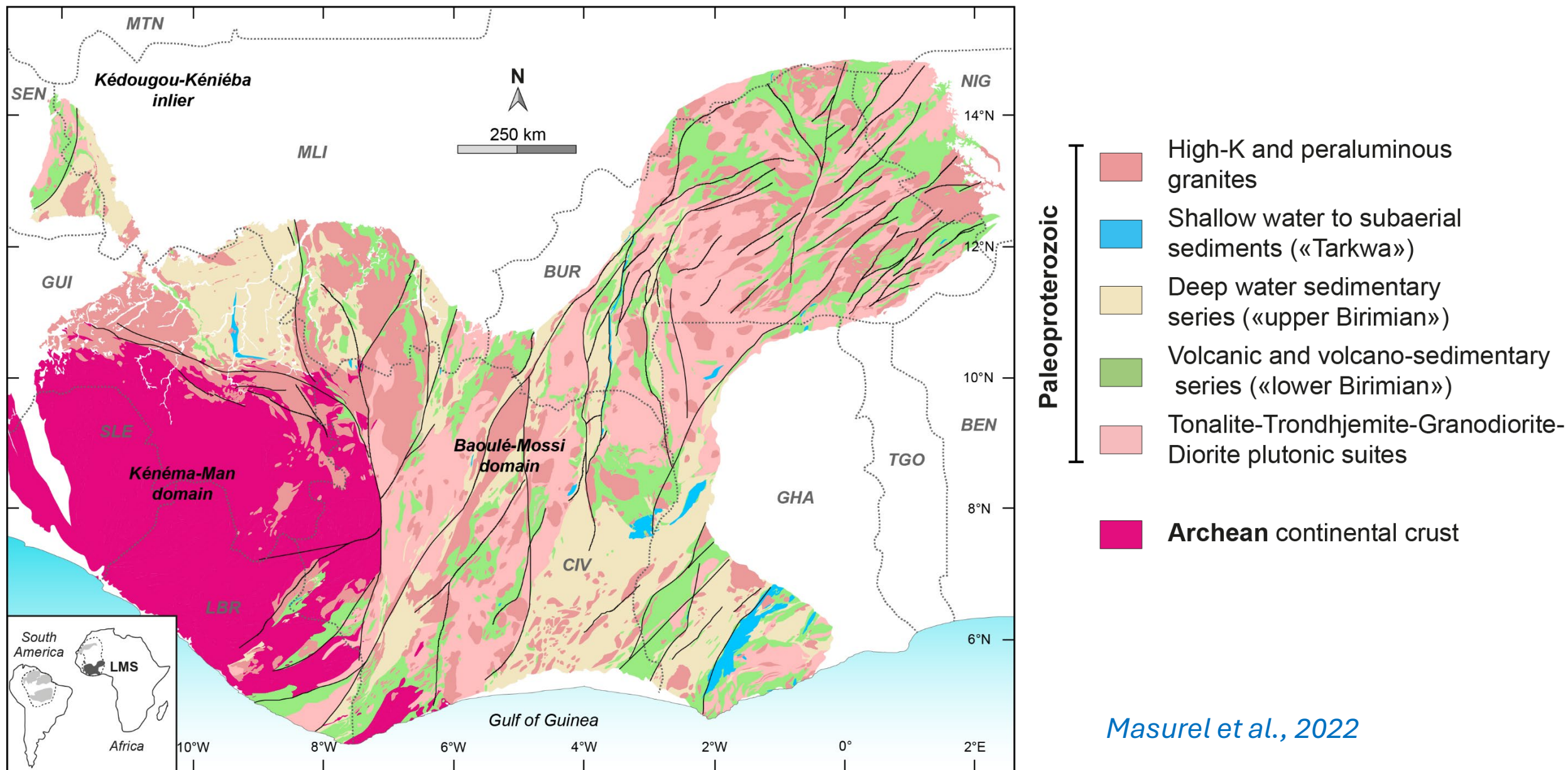


Today

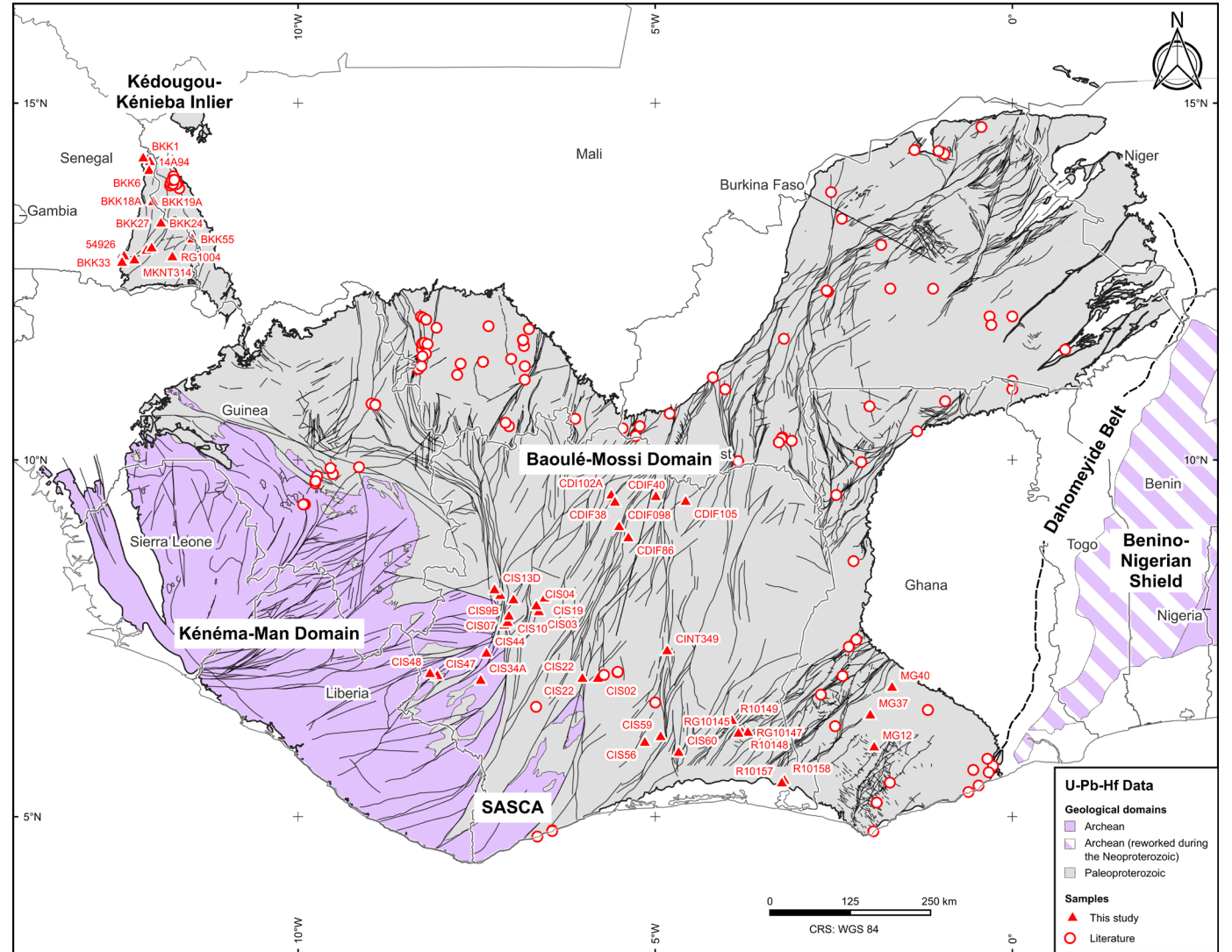
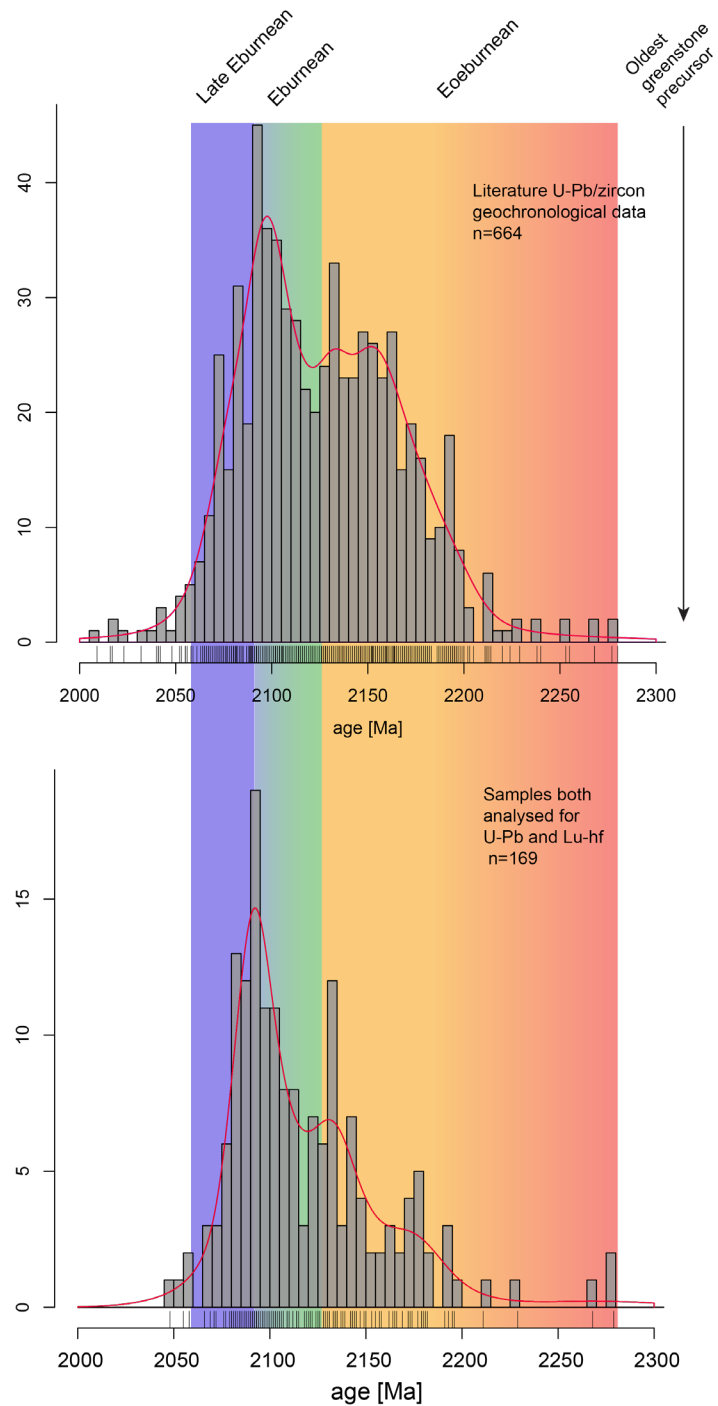
- subduction
- steady-state recycling
- no more formation of CLM
- stable



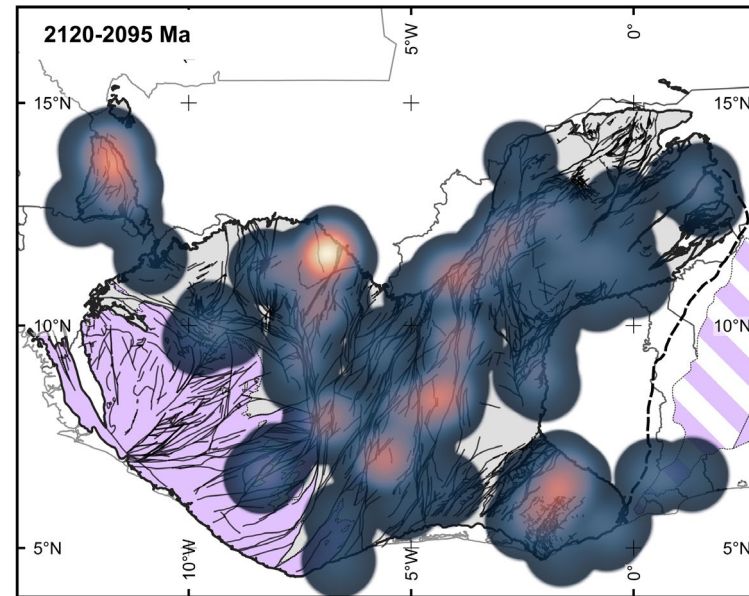
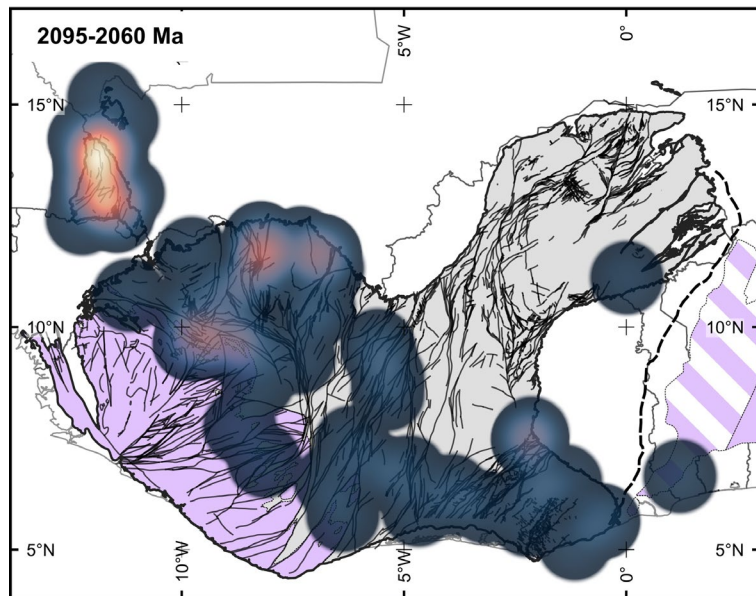
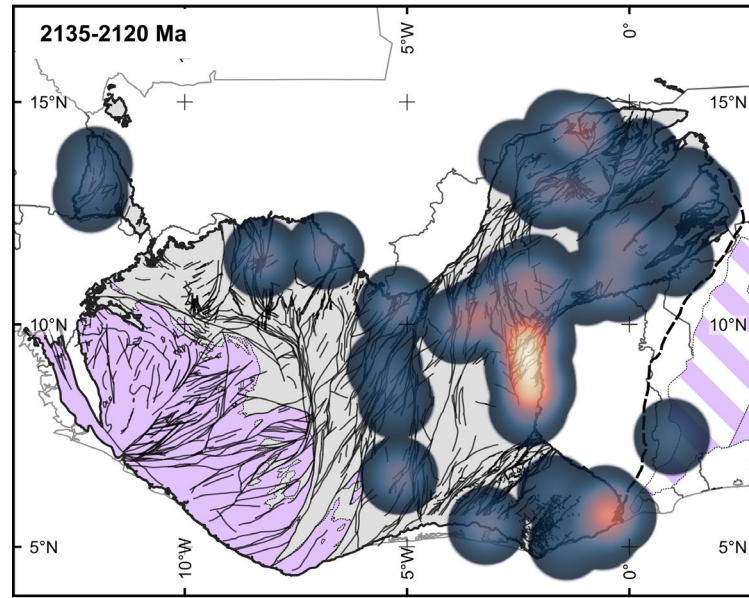
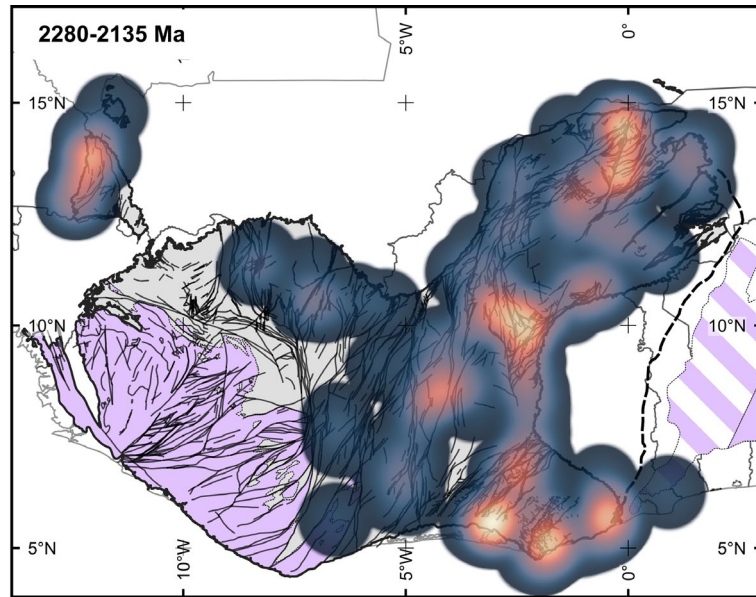
What about the south West African Craton (sWAC) ?



U-Pb felsic igneous record in the sWAC



Diachronous felsic igneous activity over the sWAC



Spatial Distribution of Magmatic Record

Geological domains

- Archean
- Archean (reworked during the Neoproterozoic)
- Paleoproterozoic

Heat map of magmatic activity (search radius = 100 km)

- Strong activity
- Little activity



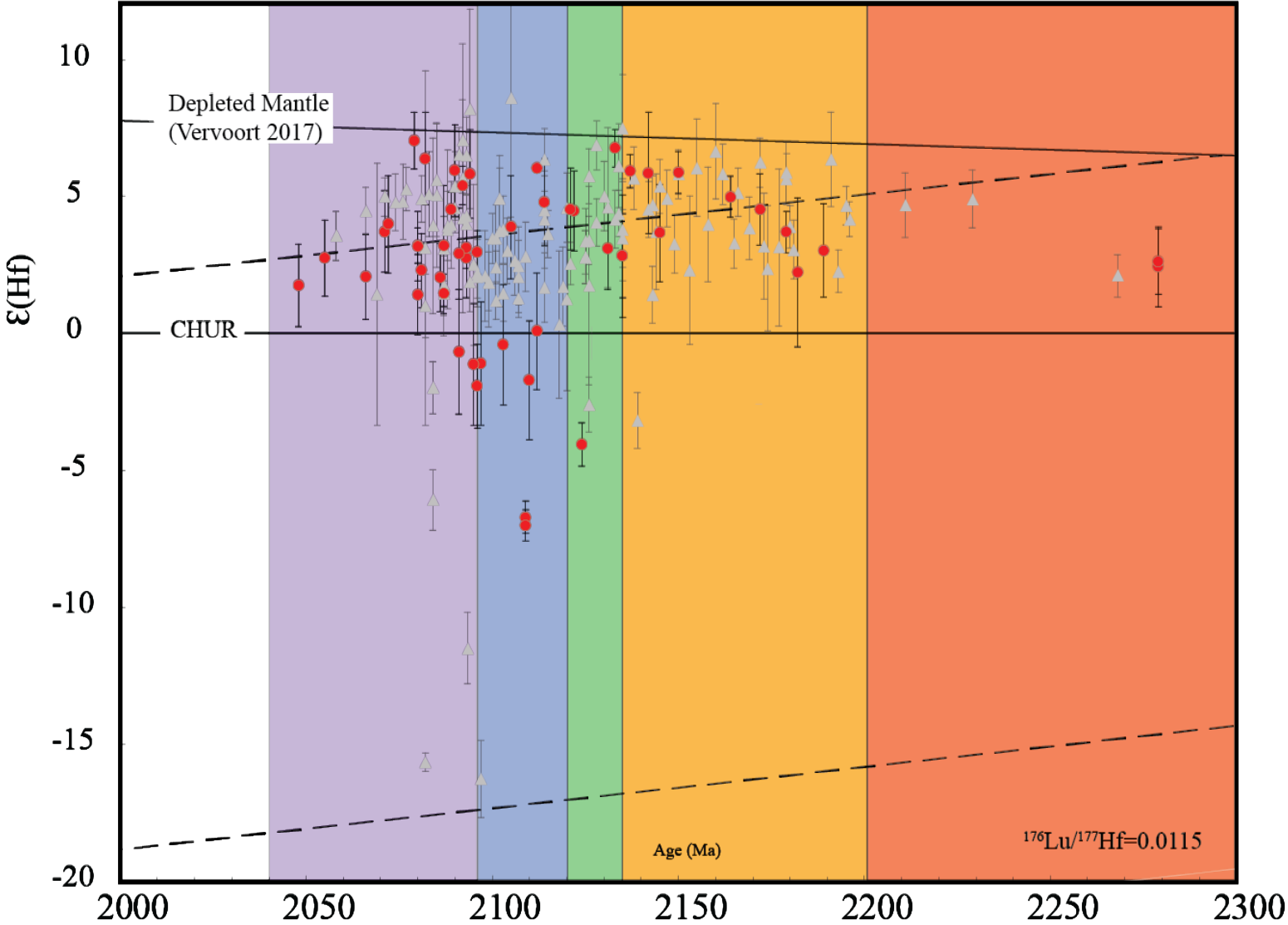
0 125 250 km

CRS: WGS 84

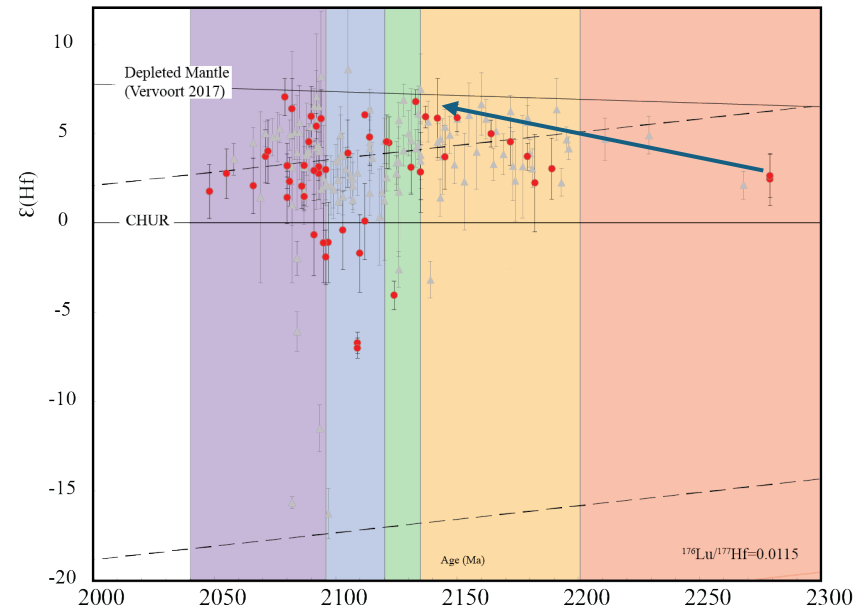
Migration westward at about 3.5 cm/yr
(Parra-Avila et al., 2017)

Thébaud et al., 2026

Lu-Hf isotopic temporal record



sWAC crustal growth and evolution during the Rhyacian

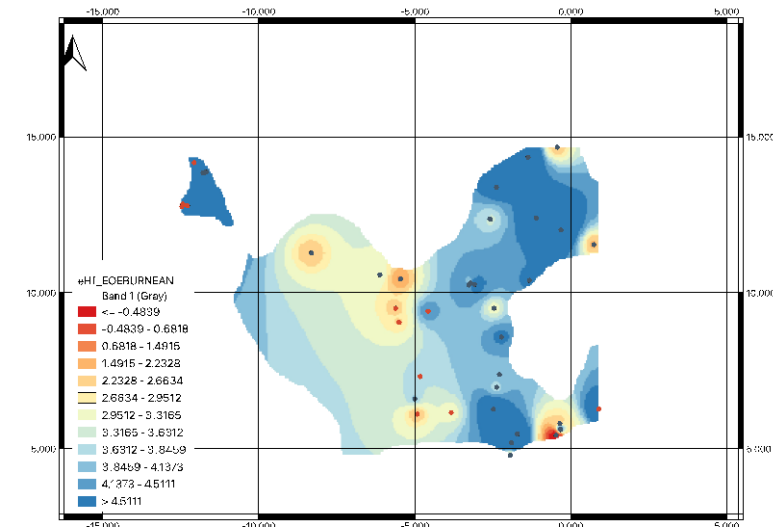


2030 to 2150 Ma

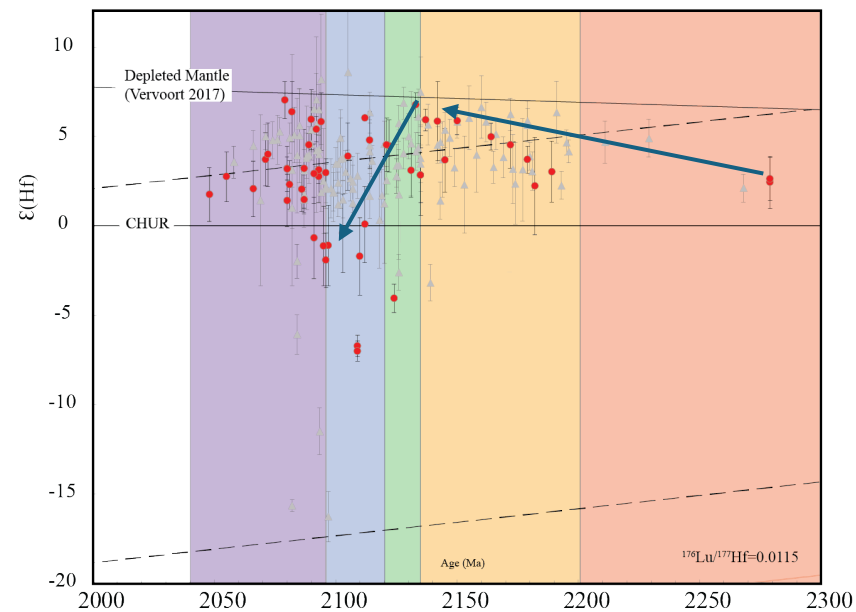
Early juvenile associated with reworking of early (2350Ma) mafic crust but also some reworking of pre-existing Archean crust.

Thébaud et al., 2026

Eoeburnean magmatism



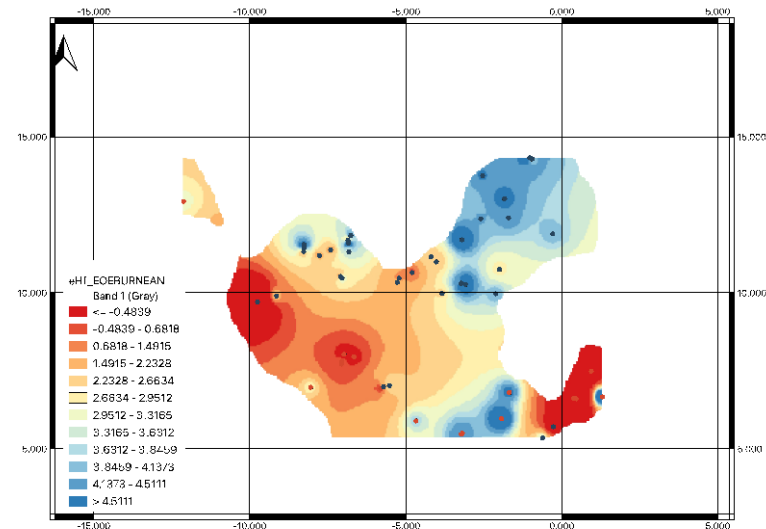
sWAC crustal growth and evolution during the Rhyacian



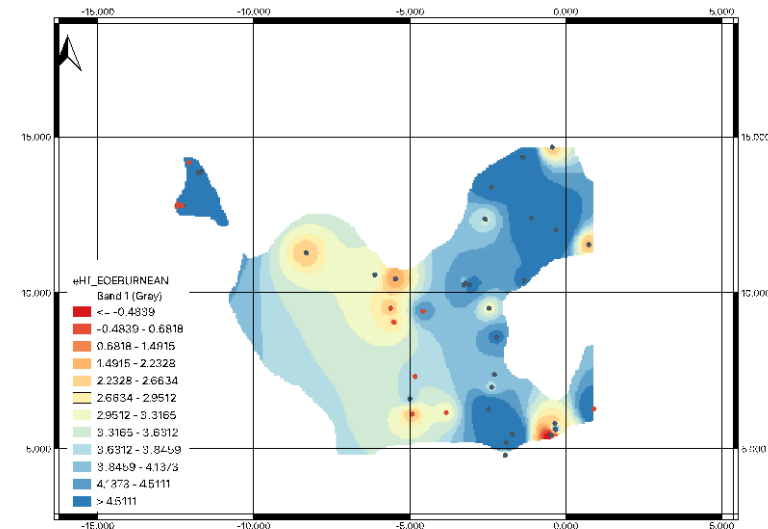
2150 to 2135 Ma
 Reworking of newly formed Eoeburnean continental crust and pre-existing Achean crust (Kenema Man Domain and BurGhana domains) in a convergent setting (Eoeburnean Orogeny).

2120 to 2095 Ma
 Ongoing reworking of crust during the Eburnean Orogeny.

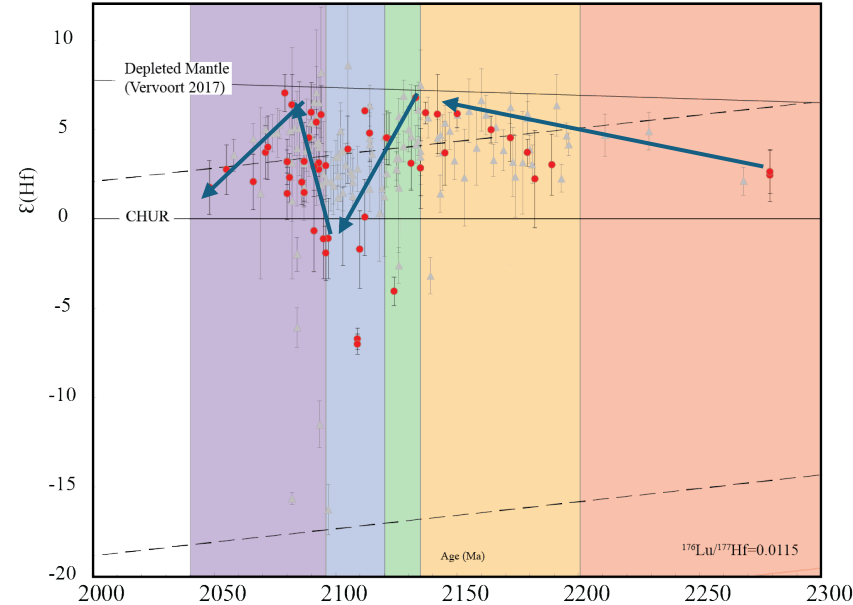
Eburnean



Eoeburnean magmatism



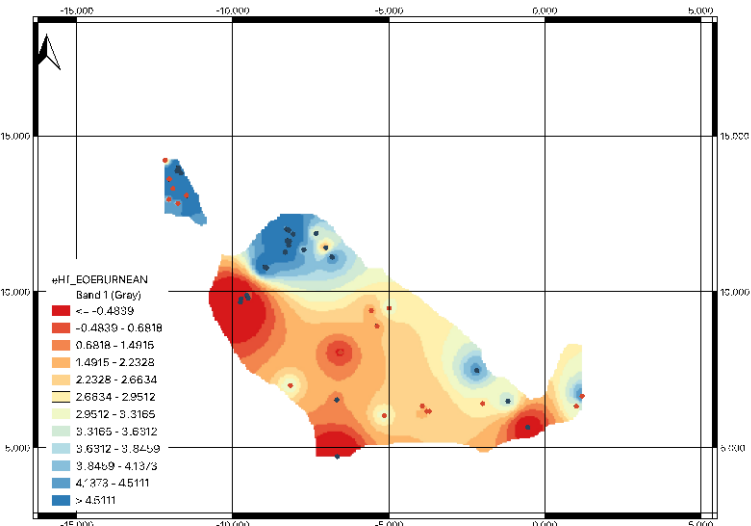
sWAC crustal growth and evolution during the Rhyacian



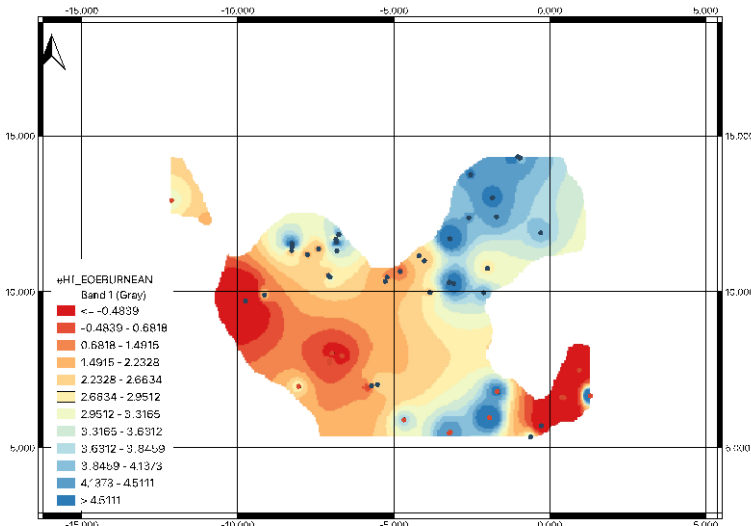
2095 to 2040 Ma
Renewed late orogenic juvenile input
(Delamination of thickened lithosphere or roll back? Eglinger et al., 2017)



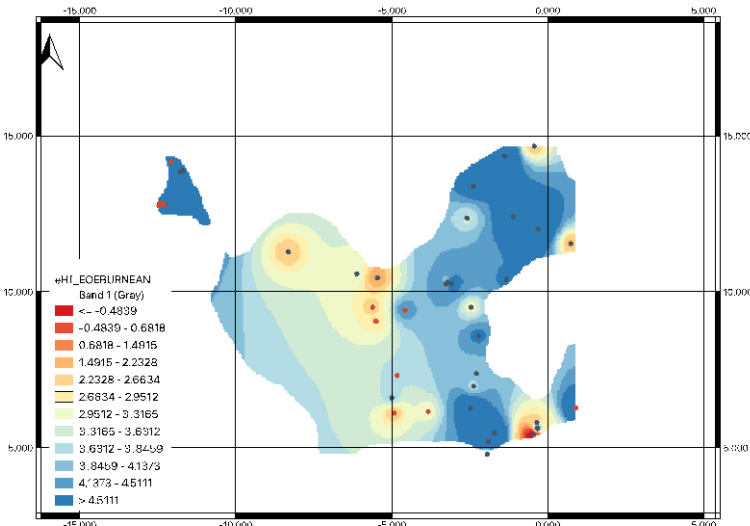
Late orogenic magmatism



Eburnean



Eoeburnean magmatism



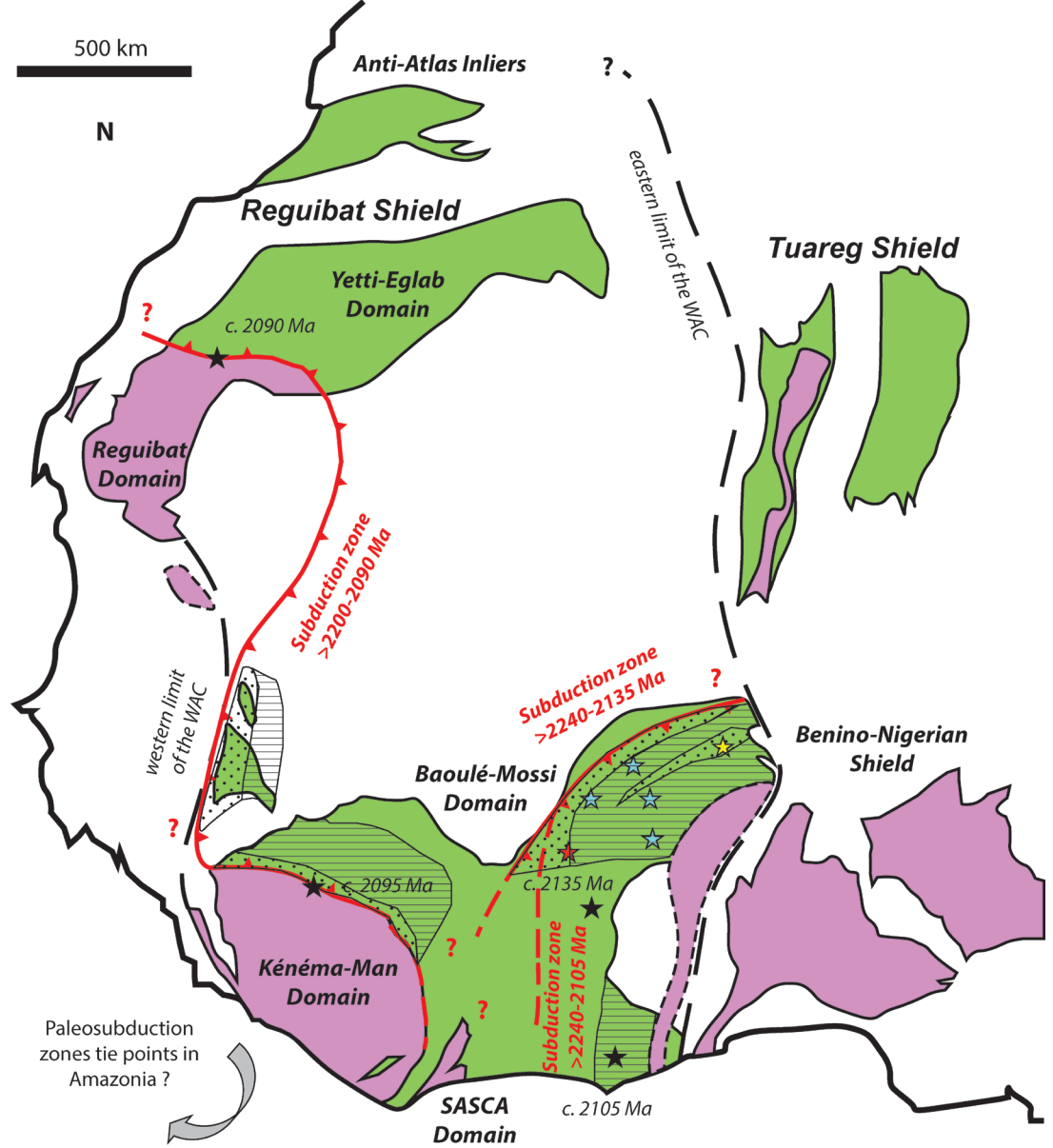
Towards a geodynamic model for the WAC

Opening of an oceanic basin at c. 2350 Ma (~3,500 km wide).

Two orogenic cycles, the Eoeburnean (~2150-2135 Ma) and the Eburnean (~2120-2095 Ma), record the progressive closure of this basin through pulses of juvenile magmatism and reworking of Archean crust.

We propose a scenario whereby two partly synchronous subduction segments developed:

- An eastern, southeast-dipping segment (c. 2240-2135 Ma) that drove Eoeburnean collision between the Baoulé-Mossi Domain and the Archean BurGhana block,
- A western, east-dipping segment (c. 2210-2070 Ma) that extended north into the Reguibat Shield and drove Eburnean collision with the Kénéma-Man Domain.



Signature ore deposits:

- ★ c. 2165 Ma porphyry Cu-Au
- ★ c. 2160 Ma porphyry Cu-Mo
- ★ c. 2175-2160 Ma VMS

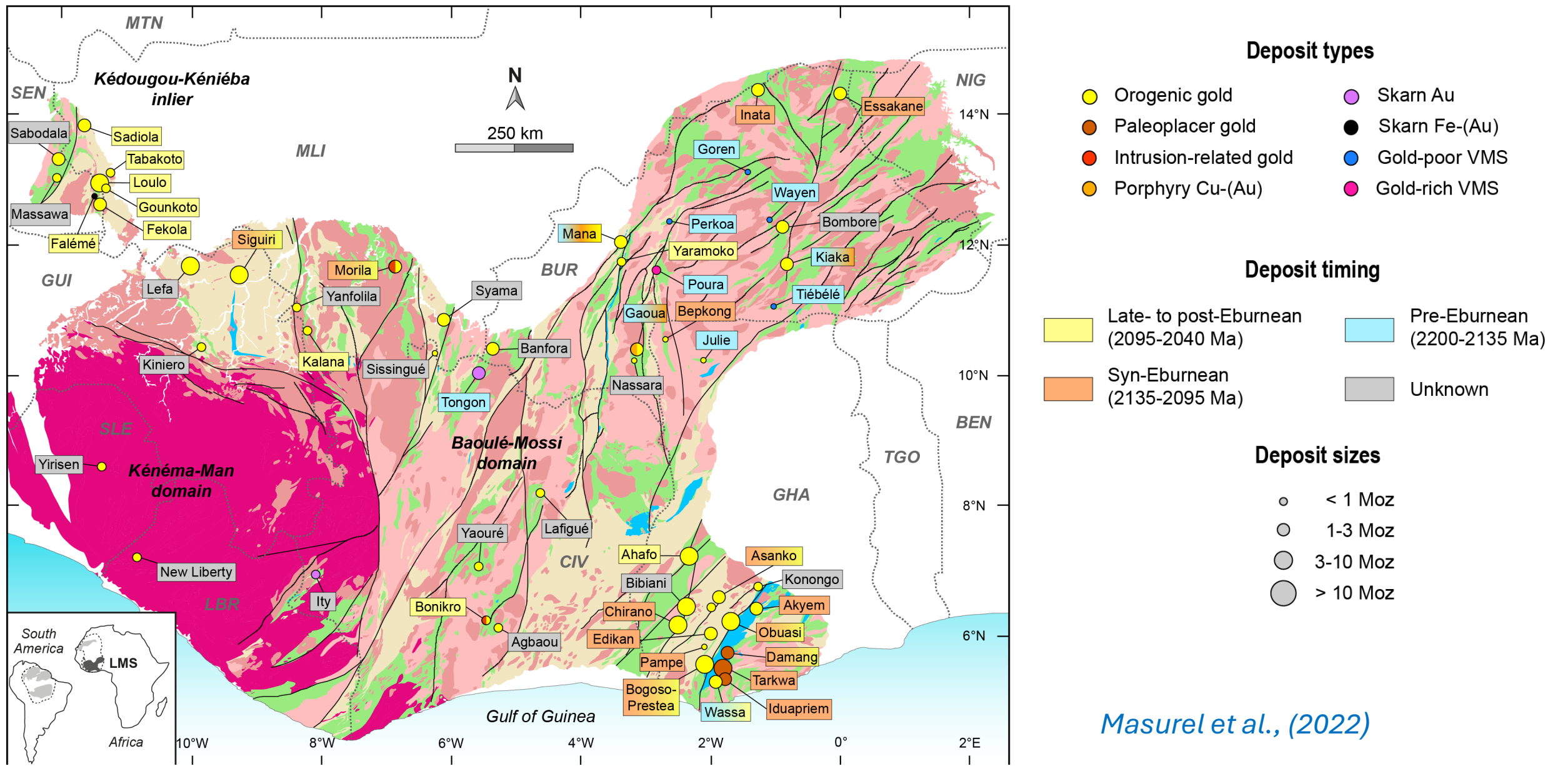
Geological domain:

- Archean
- Paleoproterozoic

Geodynamic features:

- Subduction zone (with polarity)
- Timing of arc-continent collision
- Arc
- Back Arc

South West African Craton gold endowment



A canvas to evaluate metal fertility

Au camps/districts appear to be coinciding with juvenile crustal domains (cf Yilgarn Craton, Superior).

Ni late intrusion plugs are also located at the interface or close to the interface between juvenile and evolve crustal domains.

Zn exhalative and Cu intrusion related mineralization located within the juvenile domain (arc position?).

